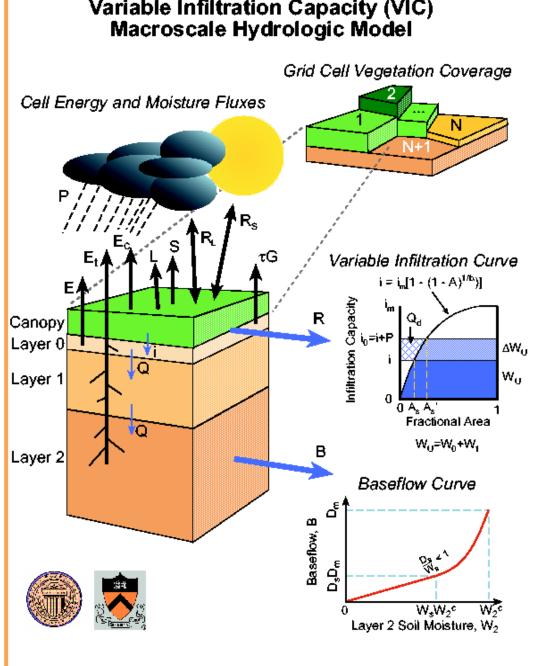
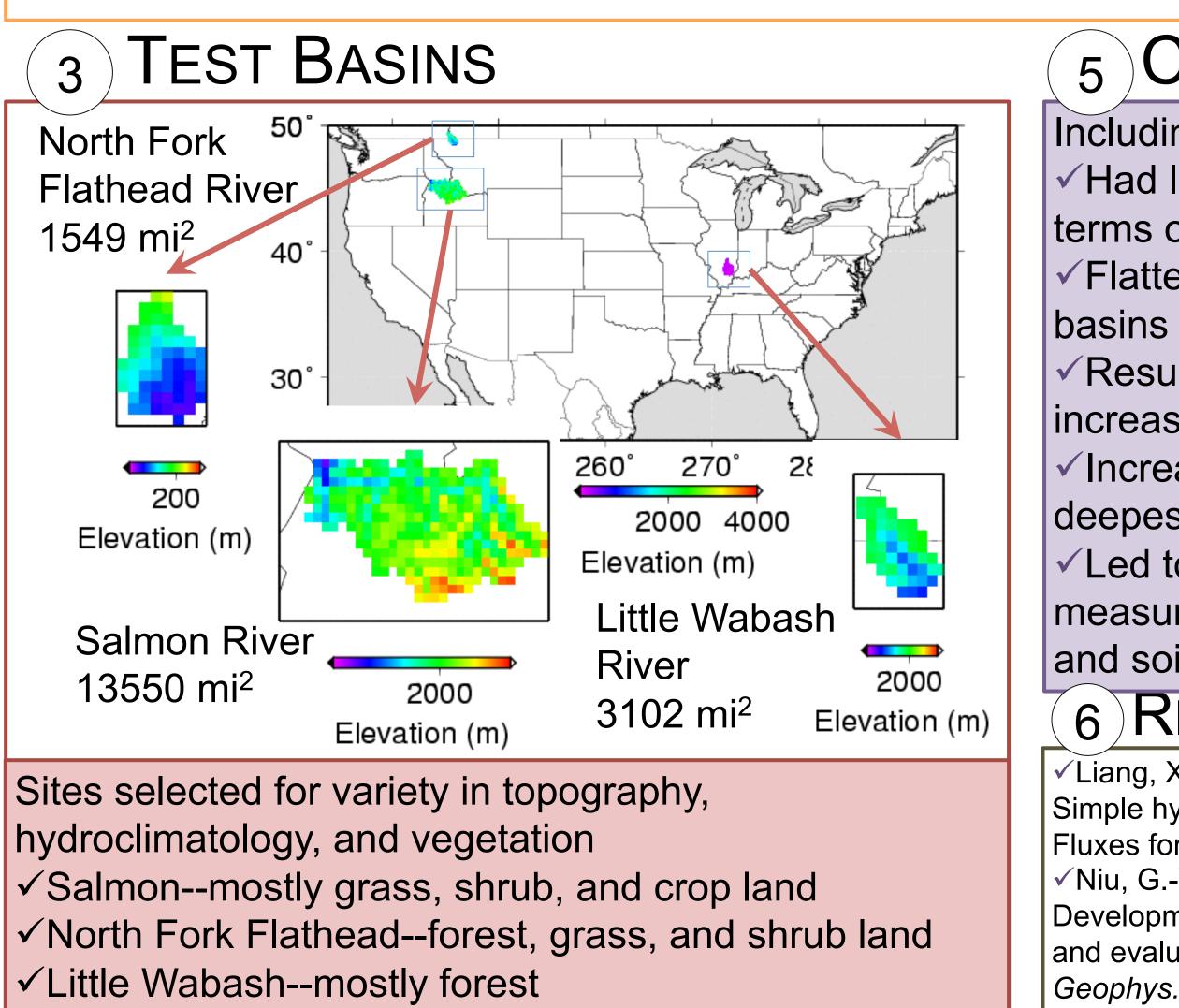
The impact of groundwater-land surface interactions on hydrologic persistence in macroscale modeling Elizabeth A. Clark and Dennis P. Lettenmaier Department of Civil and Environmental Engineering, Box 352700, University of Washington, Seattle, WA 98195 4 RESULTS Little Wabash River 1950-1999 Motivation North Fork Flathead River 1950-2000 ANNUAL Shallow groundwater interacts with the hydrologic cycle by influencing evapotranspiration and providing long-CYCLE term storage. 225 150 뉟 125 200 Precipitation \checkmark VIC and many LSMs do not explicitly incorporate shallow groundwater. 100 50 2 Evaporation 175 100 This might impact an LSM's to estimate drought duration (i.e., agricultural and hydrologic drought conditions) 50 Runoff+Baseflow 150 ge o may last longer than meteorological drought if the aquifer is drawn-down significantly; or the presence of aquifer Subsurface change 125 75 SWE change 100 storage may buffer against severe precipitation deficits). -50 50 75 É no gw -100 2) MODELING APPROACH ш -150 20 Streamflow (obs) 25 -200 -100 đ Replace baseflow component in Variable Infiltration Capacity 2 3 4 5 6 7 8 9 10 11 12 6 7 8 9 10 11 12 1 2 **Differences:** Month 1)VIC-SIMGM includes unconfined aquifer RUNOFF-250 225 -2)Subsurface flow parameterization Calibrate soil parameters (ARNO baseflow parameters,) for BASEFLOW <u></u> 200 175 5 50 PARTITIONING 150 VIC 125 Runoff (gw) Depth of lowest two soil layers and variable infiltration 100 Baseflow (gw) m 25 ш 75 Runoff (no gw) 50 W, W_sW₂¢ Calibration performed on daily flow and natural log of daily Baseflow (no gw) Layer 2 Soil Moisture, W₂ Streamflow (obs) -----10 11 12 10 11 12 $B = B_{\max} e^{-fz_{\nabla}}$ Month Month SUBSURFACE Variable Infiltration Capacity (VIC) Macroscale Hydrologic Model C with STORAGE Grid Cell Vegetation Coverage MGM Recharge Cell Energy and Moisture Fluxes Recharge Partitioning 25 - Soil layer 1 (gw) Soil layer 2 (gw) **Depth to Water** Soil layer 3 (gw) -25 Table, z_{∇} Aquifer (gw) W_a = aquifer storage $\frac{dW_a}{dW_a} = Q - B$ O = recharge to groundwater Soil layer 1 (no gw) Soil layer 2 (no gw) B = groundwater discharge Soil layer 3 (no gw) ractional Area B_{max} = maximum groundwater discharge 10 11 12 10 11 12 789 $B = B_{max}e^{-fz_{\nabla}}$ f = decay factorMonth Month K_a = hydraulic conductivity h = matric potential + gravity (elevation) potential $Q = -K_a \frac{\alpha n}{1}$ = laver depth z_{∇} = depth to water table W_sW₂^c W₂^c Layer 2 Soil Moisture, W₂ 8)Vertical soil moisture drainage STREAMFLOW $z_{\nabla} = F(z_{bot}, W_a, \text{specific yield, effective porosity})$ r = 0.36 OBS 1–month r = 0.14 OBS 3–month r = 0.09 OBS 6-month r = 0.54 OBS 1–month r = 0.26 OBS 3–month r = 0.12 OBS 6-month PERSISTENCE **TEST BASINS** 5 CONCLUSIONS Lag-correlation in Including a shallow, lumped, unconfined aquifer model: North Fork monthly streamflow ✓ Had little effect on overall model performance in r = 0.35 VIC-nL 3-month r = 0.13 VIC-nL 3-month r = 0.07 VIC-nL 6-month r = 0.11 VIC-nL 6-month with lag time of 1 r = 0.41 VIC-nL 1-month r = 0.66 VIC-nL 1-month Flathead River month on left, 3 terms of streamflow 1549 mi² months in center, and ✓ Flattened seasonal cycle in baseflow in western 6 months on right

model (VIC; Liang et al., 1994) with the SIMple Groundwater Model (SIMGM; Niu et al., 2007). VIC model with no groundwater and with groundwater (SIMGM baseflow parameters, specific yield). capacity curve exponent calibrated for both models. flow using MOCOM-UA algorithm.



9	Commonalities:	VIC
ŀ	1)Forcings: Precipitation, Tmax,	SIN
	Tmin, Wind	
	2)Sub-grid cell vegetation, roots	
e	distributed in soil layers	
	3)Surface runoff, Variable	
w _u vu	Infiltration Curve	-
	4)Soil and canopy evaporation	
	5)Transpiration from vegetation	
	6)Snow	
	7)Energy balance optional	ý



Resulted in slight summertime transpiration increases, particularly in forested basins Increased the magnitude of storage change in the deepest soil layer, particularly in western basins Led to only slightly higher lagged-correlation (as a measure of persistence) for both monthly streamflow and soil moisture, more so than observed for Q

6 REFERENCES

Liang, X., D. P. Lettenmaier, E. F. Wood, and S. J. Burges, 1994: A Simple hydrologically Based Model of Land Surface Water and Energy Fluxes for GSMs, J. Geophys. Res., 99(D7), 14,415-14,428. Niu, G.-Y., Z.-L. Yang, R.E. Dickinson, L.E. Gulden, H. Su, 2007. Development of a simple groundwater model for use in climate models and evaluation with Gravity Recovery and Climate Experiment data, J. *Geophys. Res., 112*, D07103, doi:10.1029/2006JD007522.

SUBSURFACE STORAGE PERSISTENCE

Lag-correlation in monthly total subsurface storage with lag time of 1 month on left, 3 months in center, and 6 months on right



