Macro-scale hydrology, does shallow groundwater make a difference?

Elizabeth Clark UBC/UW Meeting Sept. 25, 2009

Overview

- Introduction/Motivation
- Modeling Approaches
 - Macro-scale groundwater models
 - VIC model modifications
- Model Results
 - Little Wabash River, IL
- Conclusions

Questions

- How does the absence of shallow groundwater in a land surface model impact the water budget it produces?
- Would shallow groundwater impact drought characterization and forecasting?

Surface Water Monitor Drought Forecasts



- Forecasts runoff and soil moisture percentiles based on ensemble medians and the probability of drought conditions for 1-, 2- and 3-month lead times.
- 2 types of forecast ensembles:
 - 1960-99 climatology
 - ENSO-determined subsets from 1950-2002 climatology

Modeling approaches

Groundwater in Macro-scale Land Surface Models

• TOPMODEL-based

Increasing complexity

- Solving soil moisture for unsaturated zone and pressure head profiles for saturated zones separately
- Solving soil moisture profile by applying Richard's equation to unsaturated zone and treating water table as moving boundary
- Solving hydraulic pressure profile for unsaturated and saturated zones together

SIMGM (Niu et al. 2007)

• Solving soil moisture for unsaturated zone and pressure head profiles for saturated zones separately



VIC-ground (Liang et al. 2003)

• Solving soil moisture profile by applying Richard's equation to unsaturated zone and treating water table as moving boundary





Results for Tulpehocken Creek watershed (in Pennsylvania) with a drainage area of 172 km²

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500 05/01/92 05/01/93 05/01/94 05/01/95 05/01/96 05/01/97 05/01/98 Days (Oct. 1,1991–Sept. 30,1998)

Results for Tulpehocken Creek watershed (in Pennsylvania) with a drainage area of 172 km²

Must be computationally officient

• Must be computationally efficient

Resolution	Continental U.S.	West-wide Region
	Number of grid cells	
		Number of grid cells
1/2° x 1/2°	3,322	~2,672
1/8° x 1/8°	56,335	42,767
100 m x 100 m	~563,350,000	~427,670,000



Model Set-up

Commonalities:

- 1) Forcings: Precipitation, Tmax, Tmin, Wind
- 2) Sub-grid cell vegetation, roots distributed in soil layers
- 3) Surface runoff, Variable Infiltration Curve
- 4) Soil and canopy evaporation
- 5) Transpiration from vegetation
- 6) Snow
- 7) Energy balance optional
- 8) Vertical soil moisture drainage

Differences:

- 1) VIC-SIMGM includes unconfined aquifer
- 2) Subsurface flow parameterization



SIMple Groundwater Model

$$\frac{dW_a}{dt} = Q - R_{sb}$$
$$R_{sb} = R_{sb,max}e^{-f\vec{x}}$$
$$Q = -K_a\frac{dh}{dz}$$

Q = recharge to groundwater R_{sb} = groundwater discharge $R_{sb,max}$ = maximum groundwater discharge f = decay factor K_a = hydraulic conductivity h = matric potential + gravity (elevation) potential z = layer depth z_{∇} = depth to water table

 W_a = aquifer storage

 $z_{\nabla} = F(z_{bot}, W_a, \text{specific yield}, \text{effective porosity})$

Model Application and Results



Daily Streamflow

1953-99	VIC-nL	VIC-SIMGM
NSE of daily Q	0.69	0.62
NSE of In (daily Q)	0.42	0.59
R ²	0.75	0.66



Daily Evapotranspiration



— VIC-nL — VIC-SIMGM

Daily Soil Moisture





Streamflow Persistence

- $Q_{residual} = (Q_{i,mo} \mu_{Q,mo}) / \sigma_{Q,mo}$
- Correlations between flow in month 1 (Q residual) with flow in month n (Offset Q residual)
- VIC-SIMGM performs similarly well as VIC-nL



Soil Moisture Persistence

- $SM_{residual} = (SM_{i,mo} \mu_{SM,mo}) / \sigma_{SM,mo}$
- VIC-SIMGM soil moisture shows a slightly stronger correlation with past soil moisture



Conclusions

For the Little Wabash River in Illinois,

- VIC-SIMGM can produce comparably reasonable streamflow estimates to those of VIC-nL
- The inclusion of groundwater primarily impacts:
 - Deep layer soil moisture
 - Summertime evapotranspiration
- VIC-SIMGM has a slightly higher lagged correlation than VIC-nL but the differences seem unlikely to have a strong impact on drought forecasting.

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SIMGM: Recharge (Q)

- Water table in "aquifer"
- Recharge from Darcy's Equation:

$$Q = -K_a \frac{-z_{\nabla} - (\psi_{bot} - z_{bot})}{z_{\nabla} - z_{bot}}$$

$$= K_a + K_a \frac{\psi_{bot}}{z_{\nabla} - z_{bot}}$$

Gravitational Capillary Rise

• Hydraulic Conductivity in Aquifer $K_a = k_{bot} \frac{(1 - e^{-f(z_{\nabla} - z_{bot})})}{f(z_{\nabla} - z_{bot})}$ Water table in soil column

 No exchange between "aquifer" and soil column

 $Q_{i} = -K_{i,\nabla} \frac{(\psi_{sat} - z_{\nabla}) - (\psi_{i} - z_{i})}{z_{\nabla} - z_{i}}$ $= K_{i,\nabla} + K_{i,\nabla} \frac{\psi_{i} - \psi_{sat}}{z_{\nabla} - z_{i}}$ Gravitational Drainage Capillary Rise = Hydraulic Conductivitybetween layers based on soil texture and water content

SIMGM: Discharge (R_{sb})

- TOPMODEL-based formulation
 - Topographic (or wetness) index: $\lambda = \ln(a/\tan\beta)$
 - a = specific catchment area, tan β = local surface topographic slope

$$R_{sb} = R_{sb,max} e^{-fz_{\nabla}}$$

- *f* can be determined by sensitivity analysis or calibration against a hydrograph recession curve
- $R_{sb,max} = \alpha K_{sat}(O) e^{-\lambda} / f$
- α K_{sat}
- Issue: Not enough high resolution (~30 m x 30 m) DEM data to calculate λ . Calibrated to 16 wells in Illinois and performed sensitivity analysis to justify applying globally.



Daily Baseflow



VIC-nL simulation
 VIC-SIMGM simulation

Little Wabash 1990 to 1995

Daily Depth to Water Table Galena Freeport **Crystal Lake** St. Charles Mt. Morris Big Bend DeKalb Fermi Cambridge Stelle Monmouth Peoria **Good Hope** Kilbourne Bondville Peri Snicarte **Bondville** Springfield Janesville Greenfield Brownstown St. Peter SWS No. 2 Olney Belleville Fairfield Sparta/Eden Rend Lake Carbondale .E. Illinois College **Dixon Springs** warm wells From late 1980s ● ICN Wells From ~1997

http://www.sws.uiuc.edu/warm/sgwdata/wells.aspx

